

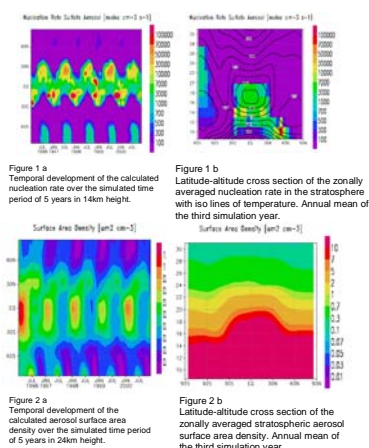
### Abstract

MIPS contributes to the understanding of the highly nonlinear interactions between chemistry, microphysics, radiation and dynamics, in studying the role of stratospheric aerosols and ice clouds in the global climate system. The investigations have been carried out with a coupled chemistry-microphysics climate model, which was developed during this project and validated for background and volcanic episodes. Different processes which were responsible for changes in the ozone concentration after the Pinatubo eruption e.g. heterogeneous chemistry, radiation, changes in the photolysis rate have been analyzed separately.

- MPI-MIPS has focused in particular on the following research activities:
- Multiannual simulations with the chemistry climate model MAECHAM4/CHEM/SAM
  - Implementation of SAM into ECHAM5
  - Investigation of aerosol chemistry interactions after the Mt. Pinatubo eruption

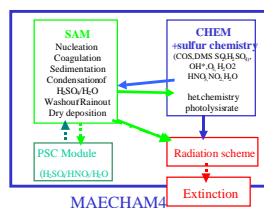
These activities are integrated in joint activities of the KODYACS partners.

### Implementation of SAM into ECHAM5



SAM has successfully implemented into ECHAM5. For the implementation in ECHAM5 SAM is rewritten in FORTRAN90 and adopted to a modular structure. A first multiannual simulation has been performed in T21 resolution (Figure 2) and is currently validated. For the near future simulations with a higher vertical and horizontal resolution are planned. In addition a coupling with the PSC module from MPI-C is intended for the summer until then both modules should run independently.

### Multiannual simulation with the chemistry climate model MAECHAM4/CHEM/SAM



We have coupled the global stratospheric aerosol model SAM (Timmreck *et al.*, 2001), which explicitly calculates microphysical processes (nucleation, condensation, coagulation, sedimentation) with the MAECHAM4/CHEM (Steil *et al.*, 2003) see KODYACS poster by Brühl *et al.* In this context SAM has been substantially revised. A new parameterisation of the homogeneous nucleation rate has been implemented (Vehkamäki *et al.*, 2002). In addition we introduced COS and H<sub>2</sub>SO<sub>4</sub> as new tracers in the model and together with MPI-C CHEM has been extended with a stratospheric sulfur scheme.

A first multiannual run has been performed with the chemistry climate model MAECHAM4/CHEM/SAM. In the upper troposphere highest particle concentration are found in NH midlatitudes in spring. The concentrations differ between land and ocean up to two orders of magnitude. The geographical distribution is similar as observed by CARIBIC (Hermann *et al.*, 2003) with high particle number concentration in the UT in summer over the Arabian Sea and at midlatitudes over Europe and lower values in the subtropics over the Middle East. In winter the model results show also a decrease over Europe in the UT towards higher latitudes. The values are however higher as observed in particular over Europe because an anthropogenic SO<sub>2</sub> emission data set from the mid 80ties is used. In the lower stratosphere the maximum number concentration is found in the tropics and is strongly linked to homogeneous nucleation.

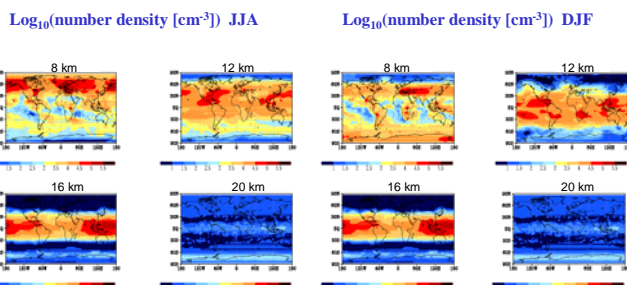


Figure 3: Decadal logarithm of particle number concentration in JJA and DJF. The model results represent a five year average of a multiannual simulation with the MAECHAM4/CHEM/SAM.

### Pinatubo effect on cirrus clouds

Different sensitivity studies have been performed with the ECHAM general circulation model for 2 1/2 years after the Mt. Pinatubo eruption (July 1991 to December 1993) to study the possible potential of volcanic emission to alter cirrus cloud properties (Lohmann *et al.*, 2003). Homogeneous cirrus formation caused by sedimenting sulphate particles produced in the eruption plume has been compared with homogeneous cirrus formation in the undisturbed atmosphere. The results turned out to be very sensitive to the assumed size distribution. While there is no trend on cloud microphysical or optical properties in our second scenario (trimodal size distribution), the first scenario (monomodal size distribution) shows a pronounced increase in ice water path and a noticeable impact on cloud radiative forcing.

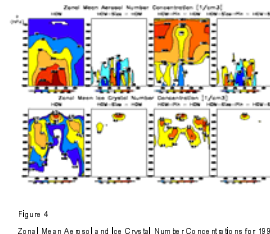


Figure 4: Zonal Mean Aerosol and Ice Crystal Number Concentrations for 1992.

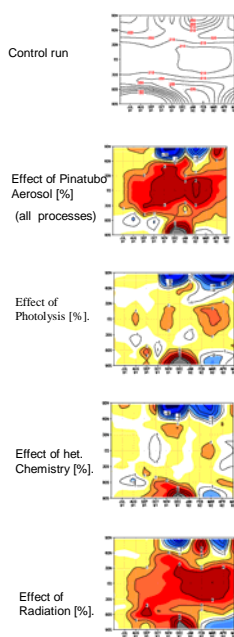
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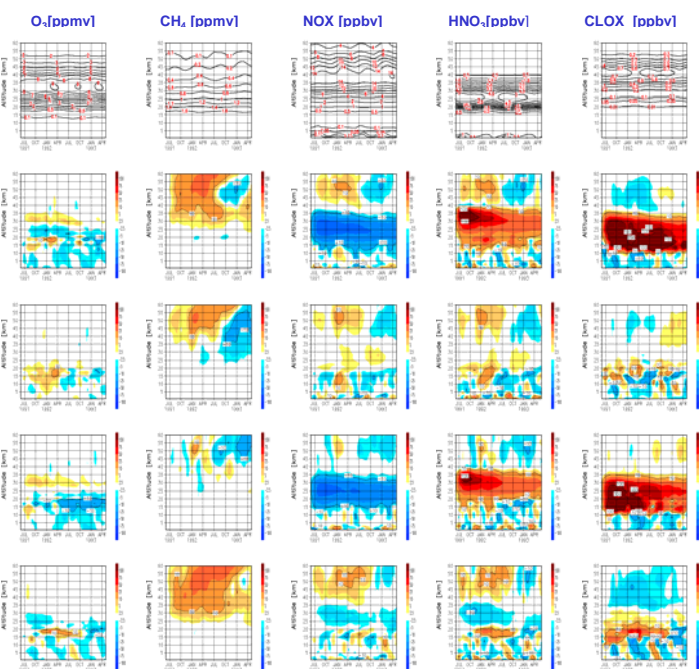
### AEROSOL CHEMISTRY INTERACTIONS IN THE STRATOSPHERE: A PINATUBO CASE STUDY

Major volcanic eruptions have a significant impact on stratospheric and tropospheric climate, chemical composition and circulation. Changes in the atmospheric trace gas concentration are a combined effect of heterogeneous chemistry, and of perturbations in the heating rates and in the photolysis rates. We have used the chemistry climate model with **interactive stratospheric chemistry and prognostic and interactive volcanic aerosol** to analyse these processes for the 1991 volcanic eruption of Mt. Pinatubo (Timmreck *et al.*, 2003).

#### Temperature changes at 50 hPa



#### Changes of Chemical Concentration (Tropical Average (30°S-30°N))



The temperature anomaly at 50 hPa is largely due to the aerosol induced radiative heating (direct effect). Heterogeneous chemistry on the volcanic aerosols particles leads to a temperature decrease of up to -1 K in October 1991 while aerosol induced changes in the photolysis rates have an opposite effect. The changes in the chemical concentration due to the volcanic aerosol are a combined effect of changes in the photolysis rates, in the heterogeneous chemistry and the heating rates. The change in the photolysis rates caused by an increase in the stratospheric particle concentration leads to two effects: a direct effect (chemistry) which is largest below the aerosol cloud and indirect effect (transport) which leads to enhanced upward transport in the first year after the eruption. Heterogeneous chemistry plays the dominant role in the aerosol containing layers between 20-30 km with increases in the ClO<sub>x</sub> concentration up to 100% and decreases in the NO<sub>x</sub> concentration of more than 50%. Aerosol induced heating leads to an uplifting of the trace gases. This results in a decrease in the aerosol containing layers and an increase in the upper stratosphere, which can clearly be seen in the NO<sub>x</sub> concentration. The tropical O<sub>3</sub> concentration decreases below 30 km due to heterogeneous chemistry and upward transport, and increases above 30 km due to a decrease in NO<sub>x</sub>.